A generic approach to explicit simulation of uncertainty in the NEMO ocean model

Jean-Michel Brankart Florent Garnier, Pierre Brasseur

Institut des Geosciences de l'Environnement Equipe de Modélisation des Ecoulements Océaniques Multi-échelles

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Since important decisions must rely on simulations, it is essential that its validity be tested, and that its advocates be able to describe the level of authentic representation which they achieved.

Summer Computer Simulation Conference (1975), cited by Richard Hamming (1997)

Motivations for a probabilistic approach

The deterministic approach is not always sufficient to describe the dynamical behaviour of the system

Comparison between simulations and observations is easier with the probabilistic approach

A good knowledge of model accuracy is necessary to solve data assimilation problems

Outline

1 Introduction

2 **Explicit simulation of uncertainties**

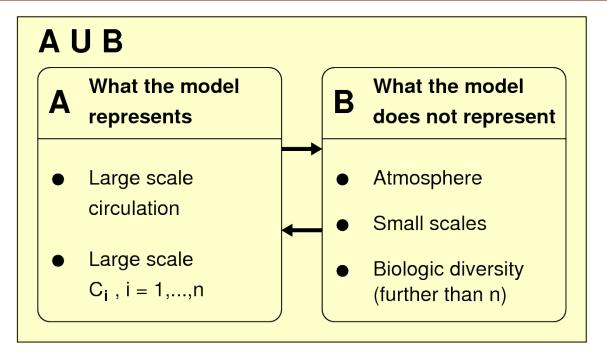
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Stochastic circulation model

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Introduction

Sources of uncertainties in ocean models



- •Even if the dynamics of **U** can be assumed deterministic, the system **A** alone **cannot be assumed deterministic**.
- To obtain a deterministic model for A, one must assumed, either that B is known (→ atmospheric forcing), or that the effect of B can be parameterized (→ paramétrisation of unresolved scales or unresolved biologic diversity).
 - → B is the main source of uncertainty in the model.

Uncertainty, as a key component of our systems

What are the uncertain components of our systems?

How to describe uncertainties?

How does it participate to the solution of inverse problems?

2

Explicit simulation of uncertainties

A first simple implementation based on autoregressive processes (1)

<u>Method</u>: explicitly simulate uncertainties in the model using *random numbers*

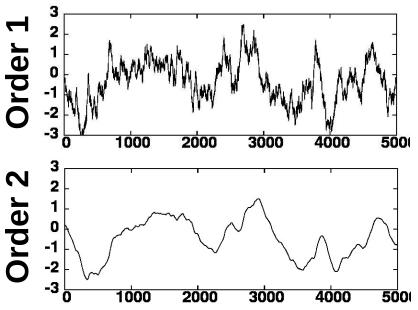
At every model grid point (in 2D or 3D), generate a set of independent Gaussian autoregressive processes:

$$\xi(t_k) = a \, \xi(t_{k-1}) + b \, w + c$$

where \mathbf{w} is a Gaussian white noise (\rightarrow order 1 process) or an autoregressive process of order n-1 (\rightarrow order n process)

Parameters *a*, *b*, *c* to specify:

mean, standard deviation and correlation timescale



A first simple implementation based on autoregressive processes (2)

Introduce a spatial correlation structure

by applying a spatial filter to the map of autoregressive processes:

$$\tilde{\boldsymbol{\xi}} = \mathcal{F}[\boldsymbol{\xi}]$$
 (filtering operator)

$$\mathcal{L}[\tilde{\boldsymbol{\xi}}] = \boldsymbol{\xi}$$
 (elliptic equation)

which can easily be made flow dependent if needed

Modify the marginal probability distributions

by applying anamorphosis transformation to every individual Gaussian variable:

$$\tilde{\boldsymbol{\xi}} = \mathcal{T}[\boldsymbol{\xi}]$$
 (nonlinear function)

for instance to transform the Gaussian variables into lognormal or gamma variables if positive noise is needed

→ This provides a generic technical way of implementing a wide range of stochastic parameterizations

Technological approach: a stochastic module in NEMO

These processes are generated using a **new module in NEMO**, and **can be used in any component** of the model (Brankart et al., 2015): circulation model, ecosystem model, sea ice model

Algorithm 1 sto_par

```
for all (map i = 1, ..., m of autoregressive processes) do
   Save map from previous time step: \xi_- \leftarrow \xi_i
   if (process order is equal to 1) then
       Draw new map of random numbers w from \mathcal{N}(0,1):
       \xi_i \leftarrow w
       Apply spatial filtering operator \mathcal{F}_i to \xi_i: \xi_i \leftarrow \mathcal{F}_i[\xi_i]
       Apply precomputed factor f_i to keep SD equal to 1:
      \xi_i \leftarrow f_i \times \xi_i
   else
       Use previous process (one order lower) instead of white
      noise: \xi_i \leftarrow \xi_{i-1}
   end if
   Multiply by parameter b_i and add parameter c_i: \xi_i \leftarrow b_i \times
   \xi_i + c_i
   Update map of autoregressive processes: \xi_i \leftarrow a_i \times \xi_- + \xi_i
end for
```

- → Generic and flexible technological approach
 - → Model independent implementation
- → Possible to simulate many kinds of uncertainty

Algorithm 2 sto par init

end for

end if

if (restart file) then

initial seed)

```
Initialize number of maps of autoregressive processes to 0:
m \leftarrow 0
for all (stochastic parameterization k = 1, ..., p) do
   Set m_k, the number of maps of autoregressive processes re-
   quired for this parameterization
   Increase m by m_k times the process order o_k: m \leftarrow m +
   m_k \times o_k
end for
for all (map i = 1, ..., m of autoregressive processes) do
   Set order of autoregressive processes
   Set mean (\mu_i), standard deviation (\sigma_i) and correlation
   timescale (\tau_i) of autoregressive processes
   Compute parameters a_i, b_i, c_i as a function of \mu_i, \sigma_i, \tau_i
   Define filtering operator \mathcal{F}_i
   Compute factor f_i as a function of \mathcal{F}_i
end for
Initialize seeds for random number generator
for all (map i = 1, ..., m of autoregressive processes) do
   Draw new map of random numbers w from \mathcal{N}(0,1): \xi_i \leftarrow
   Apply spatial filtering operator \mathcal{F}_i to \xi_i: \xi_i \leftarrow \mathcal{F}_i[\xi_i]
   Apply precomputed factor f_i to keep standard deviation
   equal to 1: \xi_i \leftarrow f_i \times \xi_i
   Initialize autoregressive processes to \mu + \sigma \times w: \xi_i \leftarrow \mu +
   \sigma \xi_i
```

Read maps of autoregressive processes and seeds for the random number generator form restart file (thus overriding the

List of uncertainties that have been implemented in NEMO using this generic stochastic module

In the circulation model

Effect of unresolved scales in the equation of state Uncertainties in the parameterized tendencies (SPPT scheme) Uncertainties in the horizontal momentum diffusion Uncertainties in vertical diffusion (lognormal distribution) Uncertainties in bulk parameters C_D , C_E , C_H (gamma distribution)

In the ecosystem model

Effect of unresolved scales in the SMS terms of the equations Uncertainties in model parameters (primary production and grazing)

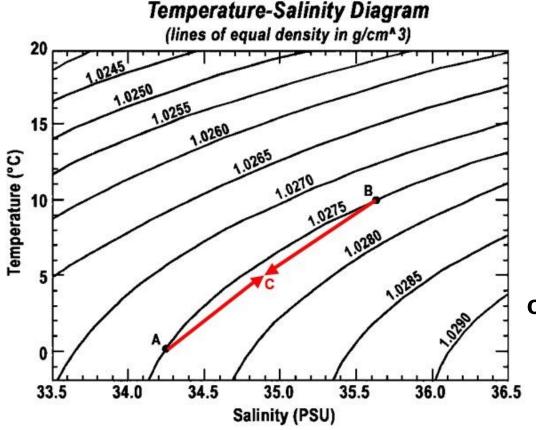
In the sea-ice model

Uncertainties in ice strength (gamma distribution)
Uncertainties in ice/atmosphere drag (gamma distribution)
Uncertainties in ice/ocean drag (gamma distribution)
Uncertainties in ice albedo (beta distribution)

Stochastic circulation model

Uncertainties in the computation of density

In the model, the large-scale density is computed form large-scale temperature and salinity, using the sea-water equation of state.



(a)
Mixing waters of equal density but different T&S systematically increases density (cabbeling)

Averaging T&S equations systematically overestimates density (in a fluctuating, non-deterministic way)

Because of the nonlinearity of the equation of state, unresolved scales produce an average effect on density.

Stochastic equation of state for the large scales

Stochastic parameterization (Brankart, 2015)

using a set of random T&S fluctuations ΔT_i et ΔS_i , i=1,...,p

to simulate unresolved T&S fluctuations

$$ho = rac{1}{p} \sum_{i=1}^p
ho \left[T + \Delta T_i, S + \Delta S_i, p_0(z)
ight] \quad ext{with} \quad \sum_{i=1}^p \delta T^{(i)} = 0 \, , \, \sum_{i=1}^p \delta S^{(i)} = 0$$

No effect if the equation of state is linear. Proportional to the square of unresolved fluctuations.

Correction $\Delta \rho$ applied in the thermal wind equation, as in the semi-prognostic method of Greatbatch et al. (2004)

No direct modification of T&S; no enhanced diapycnal mixing.
T&S only modified indirectly through a modification
of the main currents

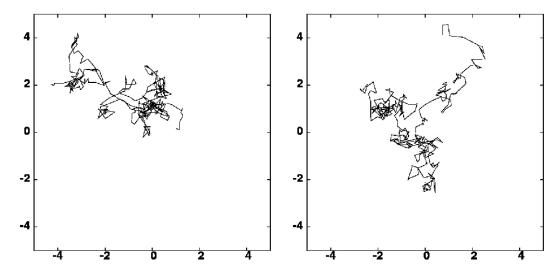
Random walks to simulate unresolved temperature and salinity fluctuations

Computation of the random fluctuations ΔT_i et ΔS_i

as a scalar product of the local gradient with random walks ξ_i

$$\Delta T_i = \boldsymbol{\xi}_i \cdot \nabla T$$
 and $\Delta S_i = \boldsymbol{\xi}_i \cdot \nabla S$

Random walks



Assumptions

AR1 random processes

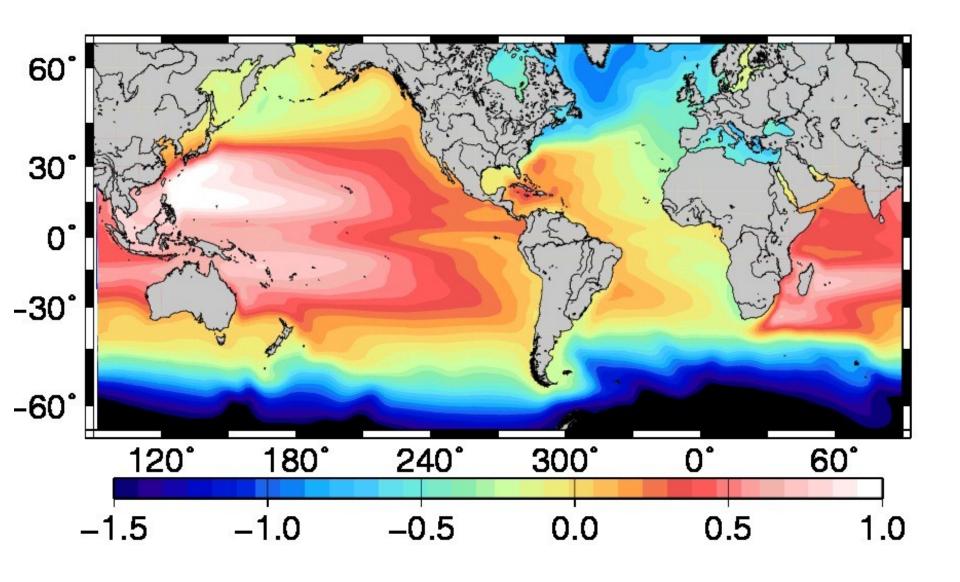
uncorrelated on the horizontal

fully correlated along the vertical

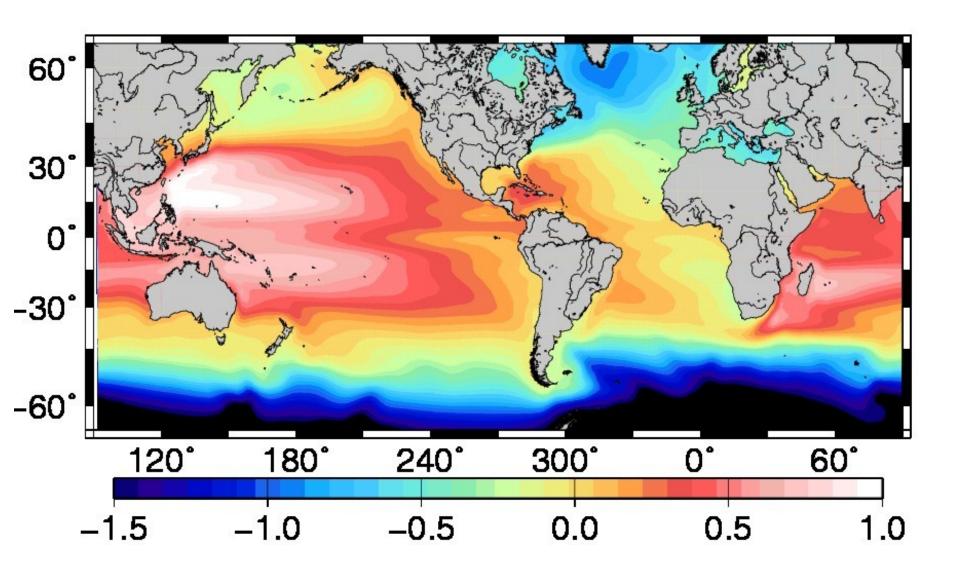
5-day time correlation

horizontal std: 2-3 grid points vertical std: <1 grid point

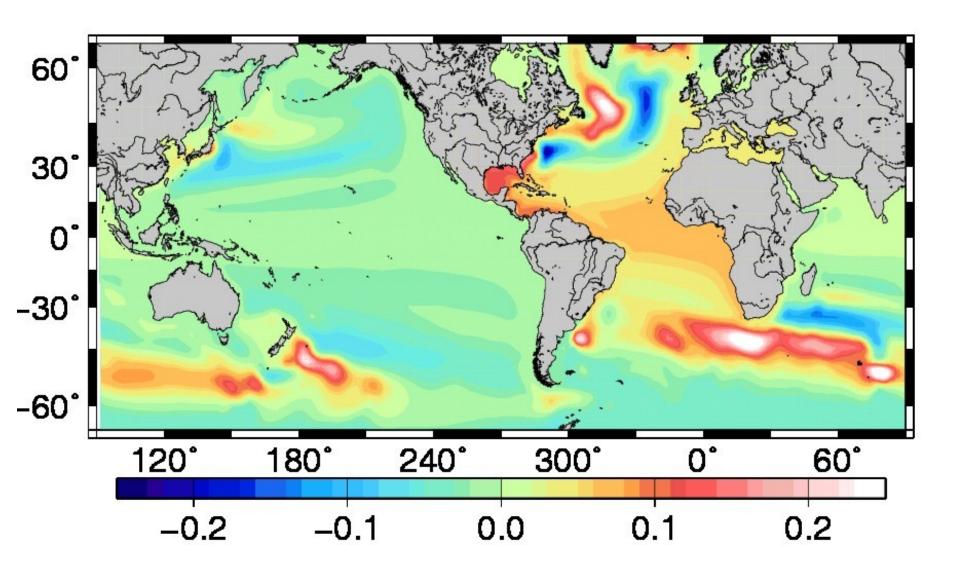
Mean sea surface elevation (standard)



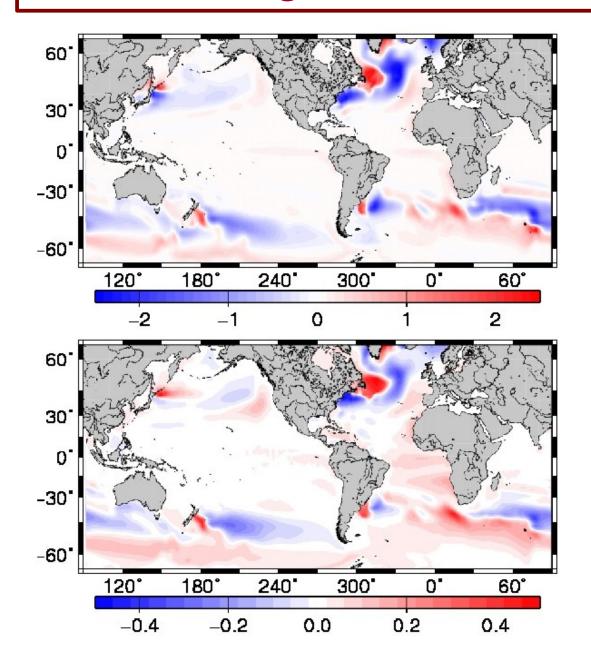
Mean sea surface elevation (stochastic)



Mean sea surface elevation difference



Averaged SST & SSS difference



Modification of the mean flow



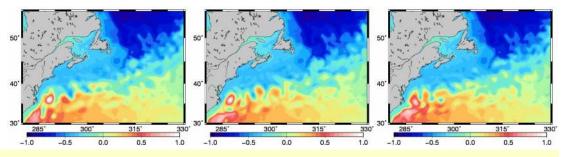
Modification of the mean SST & SSS



Modification of air/sea interactions

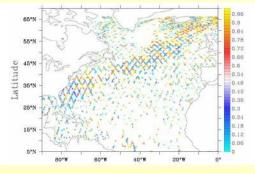
Ensemble of mesoscale flows

SSH pdf ensemble NATL025: 3 members among 96

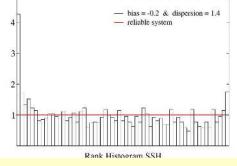


→ comparison to observation (rank histogram, CRPS score, etc.)

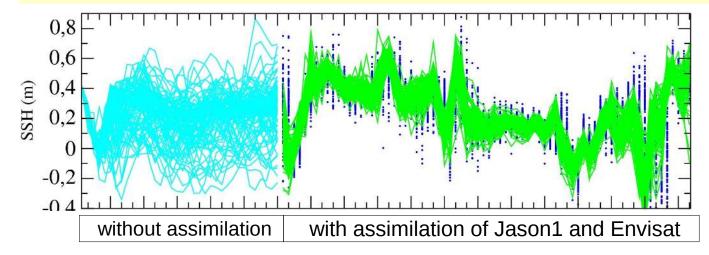
Map of ranks of JASON altimetric data



Rank histogram to check relability



→ assimilation of altimetric data (Candille et al., 2015)



Time evolution of the pdf

from June 2005 to December 2006

4

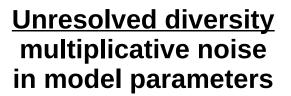
Stochastic biogeochemical model

Stochastic biogeochemical model

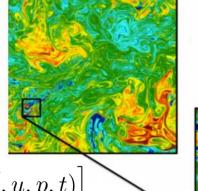
Multiple sources of uncertainty in ecosystem model: <u>unresolved biological diversity</u>, <u>unresolved scales</u>, etc.

Unresolved scales
stochastic processes
explicitly simulating
unresolved fluctuations of C_i

$$\frac{\partial C}{\partial t}|_{bio} = \frac{1}{2} \left[SMS(\overline{C} + \delta \overline{C}, u, p, t) + SMS(\overline{C} - \delta \overline{C}, u, p, t) \right]$$
$$\delta \overline{C} = \xi(t) \overline{C}$$



$$\frac{\partial C}{\partial t}|_{bio} = SMS(C, u, p', t) \quad avec \quad p' = \exp[\xi(t)]p$$



Assumptions

perturb 8 parameters governing primary production and grazing

fully correlated along the vertical

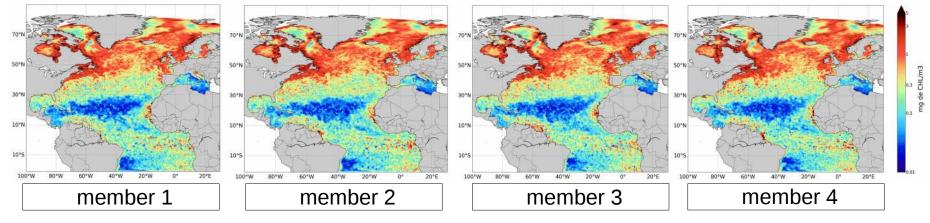
30-day time correlation

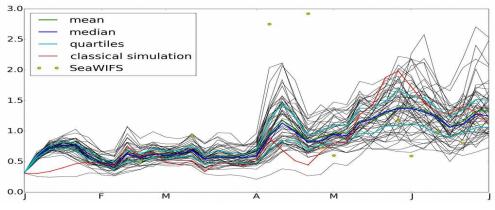
- → Considerable effect on the mean behaviour of the system
- → Increase of patchiness (→ ocean colour data)

Ensemble simulation of the ecosystem

Probability distribution of chlorophyll concentration

as simulated here by ensemble NATL025/PISCES with stochastic parameterization of uncertainty: 4 members among 50





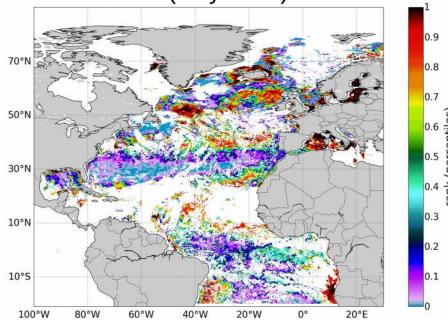
Time evolution of the pdf for phtyoplankton

from January to June 2005

→ description of biogeochemical uncertainties (Garnier et al., 2016)

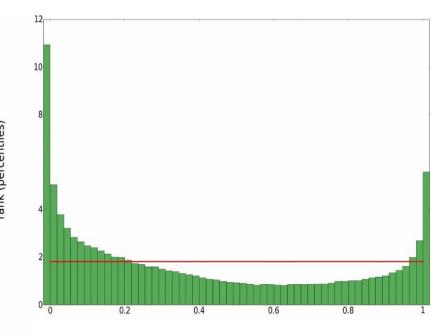
Comparison to ocean colour observations

Rank of SeaWifs observations in the ensemble simulation (May 2005)



The ensemble spread is already sufficient to include more than 80% of the observations (accounting for a 30% observation error)

Rank histogram for SeaWifs over the whole domain



The ensemble is not far from being reliable, even if still underdispersive (too many observations in the external ranges of the ensemble)

- → objectively test the consistency of simulations as compared to observations
- → prerequisite to ocean colour data assimilation

Conclusions

Probabilistic model and ensemble simulations

The NEMO model becomes <u>probabilistic</u>; il is seen as a complex system, built up from uncertain components

→ The goal of ocean modelers is then to build a model as informative as possible at the lesser cost.

This probabilistic description requires ensemble simulations

- → Objective comparison between simulations and observations
- → Deal with model uncertainty in ocean data assimilation systems

An appropriate simulation of uncertainty is necessary to make the link between model, observations, and data assimilation systems

Uncertainty is bound to become a key constituent of the systems that we are using in oceanography, not something that can be thought separately from the results

Properly dealing with uncertainty will require an integrated engineering approach at the interface between oceanography and applied mathematics

Error is viewed therefore not as an extraneous and misdirected of misdirecting accident, but as an essential part of the process under consideration.

John von Neuman (1956), in « Probabilistic logics and the synthesis of reliable organisms from unreliable components ».